

1 **Comparing different profiles to**
2 **characterize the atmosphere for three**
3 **MODIS TIR bands**

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8 **ABSTRACT**

9 Accurate Land surface temperature (LST) retrievals from sensors aboard orbiting
10 satellites are dependent on the corresponding atmospheric correction, especially in the
11 Thermal InfraRed (TIR) spectral domain (8-14 μm). To remove the atmospheric effects from at-
12 sensor measured radiance in the TIR range it is needed to characterize the atmosphere by
13 means of three specific variables; the upwelling path and the hemispherical downwelling
14 radiances plus the atmospheric transmissivity. Those variables can be derived from the
15 previous knowledge of vertical atmospheric profiles of air temperature and relative humidity
16 at different geo-potential heights and pressures.

17 In this work, the above mentioned atmospheric variables were analyzed for three
18 specific weather station site located in Spain, at three different altitudes. These variables were
19 calculated with atmospheric profiles retrieved from three different sources; The National
20 Centers for Environmental Prediction (NCEP) web-tool atmospheric profiles calculator, the
21 MODIS (MOD07) product and the radiosoundings available on the web of the University of
22 Wyoming (WYO), which are launched by the Agencia Estatal de Meteorologia (AEMET), in the
23 particular case of Spain. Atmospheric profiles from 2010 to 2013 were obtained to carry out
24 the present study.

25 Results from comparison of these three different atmospheric profiles show that the
26 NCEP profiles characterize the atmosphere in a better manner than MOD07. Average results
27 values of the three MODIS spectral bands 29, 31 and 32 show a BIAS of $0.06 \text{ Wm}^{-2}\mu\text{m}^{-1}\text{sr}^{-1}$ and
28 RMSE of $\pm 0.2 \text{ Wm}^{-2}\mu\text{m}^{-1}\text{sr}^{-1}$ for upwelling radiance, a BIAS of $0.13 \text{ Wm}^{-2}\mu\text{m}^{-1}\text{sr}^{-1}$ and RMSE of
29 $\pm 0.3 \text{ Wm}^{-2}\mu\text{m}^{-1}\text{sr}^{-1}$ for the downwelling radiance and a BIAS of -0.008 and RMSE of ± 0.03 for the
30 atmospheric transmissivity.

31 In terms of simulated LST, these errors yield a deviation of $\pm 0.9 \text{ K}$ when applying a
32 single-channel method.

33 **Keywords:** NCEP; MOD07; Radiosoundings; LST; Atmospheric profiles; MODTRAN.

34 1. INTRODUCTION

35 Surface radiance measurements taken by Remote Sensing instruments aboard
36 satellites are affected by the atmosphere. In the particular case of the thermal infrared region
37 ($3\text{-}14 \mu\text{m}$), there exists two specific spectral ranges located at $3.7\text{-}4.1 \mu\text{m}$ and $8\text{-}14 \mu\text{m}$, where
38 the atmosphere shows the lowest radiative effect, mainly due to the water vapor content (W).

39 The Land Surface Temperature (LST) is a variable of great interest in numerous
40 meteorological and climatological studies, and its accurate retrieval is of prime interest to
41 estimate energy and water fluxes budgets between the surface and atmosphere (Sánchez et
42 al., 2011 and 2014). The main concerns in retrieving LST from satellite data are the
43 atmospheric and emissivity correction (Jiang and Liu, 2014). Currently LST can be obtained
44 from several algorithms dependent on the sensor specifications (Zhou et al., 2012). Two of
45 these techniques are the single-channel (SC, Vlassova et al., 2014) and the Temperature and
46 Emissivity Separation (TES, Hulley et al., 2014) methods. Both algorithms derive the LST from
47 inversion of the radiative transfer equation (RTE), which at the same time needs the previous
48 knowledge of a characterized atmosphere, among other factors. The radiation measured with

49 a radiometric sensor is composed of a double contribution: first, the radiation directly emitted
 50 by the surface, and secondly, the radiation reflected in the surface, coming from the
 51 surroundings and the atmosphere. In addition, if the measurement is taken from satellite, the
 52 atmosphere contributes in two different ways: on the one hand, absorbing part of the surface
 53 radiation and in the other, emitting radiation directly to the sensor. RTE connects the at-sensor
 54 radiation and the radiation emitted by the surface through an energy balance defined as:

$$55 \quad L_i^{TOA} = \tau_i [\varepsilon_i B_i(T) + (1 - \varepsilon_i) L_{i,hem}^\downarrow] + L_i^\uparrow(\theta) \quad (1)$$

56 where τ_i is the atmospheric transmissivity at the spectral range i , ε_i is the surface spectral
 57 emissivity, $B_i(T)$ is the Planck function for black body spectral radiance at temperature T ,
 58 $L_{i,hem}^\downarrow$ is the downwelling hemispheric radiance, $L_i^\uparrow(\theta)$ is the upwelling path radiance at zenith
 59 angle θ , and L_i^{TOA} is the radiance measured by the sensor at the top of the atmosphere. All the
 60 surface and atmospheric factors exposed in Eq. (1) can be found explicitly defined in the
 61 reviewing publication of García-Santos et al. (2010). The W present in the atmosphere affects
 62 directly to the value of these atmospheric parameters ($\tau_i, L_{i,hem}^\downarrow, L_i^\uparrow(\theta)$).

63 Nowadays, the most suitable way to characterize the atmosphere is by using a
 64 radiative transfer code (RTC, Berk et al., 2011), which calculates the atmospheric factors from
 65 introducing vertical profiles of pressure, air temperature and humidity at different levels of
 66 altitude. Probably, the best representative profile of the corresponding atmosphere is
 67 obtained from radiosounding data, acquired with a launched balloon. However, this data is
 68 rarely available at the time and location of the measurements acquisition. As an alternative,
 69 there exists the possibility to obtain an atmospheric profile derived from the spectral features
 70 of overpassing satellite sensor as well as with interpolating models (in space and time) based
 71 on radiosounding data acquired close to a selected point (Jimenez-Muñoz et al., 2010).

72 The objective of this paper is to analyze differences on the atmospheric variables, τ_i ,
 73 $L_{i,hem}^\downarrow, L_i^\uparrow(\theta)$, after applying on MODTRAN RTC (Berk et al., 2011) different vertical profiles of

74 cloudless days. These profiles are obtained from three different sources: a) Modeled MODIS
75 spectral measured radiances (MOD07, Borbas et al., 2011) profiles. b) Spatial and temporal
76 interpolated National Centers for Environmental Prediction (NCEP) atmospheric profiles (Barsi
77 et al., 2005). c) Radiosounding data measured by balloons launched by AEMET and showed on
78 the web of the Department of Atmospheric Science in the University of Wyoming (WYO).

79 In previous studies carried out by Coll et al. (2012) and Li et al. (2013), the LST obtained
80 from satellite data using SC methods were compared with in situ ground LST measurements,
81 after applying different atmospheric profiles obtained from the NCEP and the MOD07 product.
82 In both cases, these studies obtained good results for the LST comparison, but MOD07
83 introduced greater errors in the LST retrievals (± 1.0 K for NCEP and MOD07 in Coll et al. 2012
84 and ± 1.1 K for NCEP and ± 1.2 K for MOD07 in Li et al. 2013). This paper faces the same
85 objective pursued by Coll et al. (2012) and Li et al., (2013) but in a different and profuse
86 manner. The most important contribution in this paper is that the comparison of atmospheric
87 parameters obtained from NCEP and MOD07 profiles with in situ radiosounding data has been
88 done for a period of 4 years (from 2010 to 2013), and for three different sites with differences
89 in the heights above sea level as well as in the distances from the corners used by NCEP to
90 interpolate the profiles, yielding to more representative statistical results. In addition, results
91 and statistics from MODIS band 29 ($8.4 - 8.7 \mu\text{m}$) are included. This band, which is not
92 analyzed before in the other papers, is added to the study. It is important an accurate
93 atmospheric correction between $8.4-8.7 \mu\text{m}$, which is required in mineralogical and geological
94 research purposes. For instance, mapping geological presence of minerals, like quartz, based
95 on measurements in the reststrahlen region $8-9 \mu\text{m}$, where the emissivity of quartz decreases
96 in a very pronounced manner.

97 Section 2 describes the methodology applied to reach the fixed objective. Section 3,
98 shows the results obtained and the corresponding discussions. A simulation study was carried
99 out in section 4 to evaluate the effect in terms of LST when applying SC method with NCEP

100 and MOD07 profiles if in-situ radiosounding data is not available. Finally, the main conclusions
101 of the study are given in section 5.

102 **2. MATERIALS AND METHODS**

103 **2.1 Site**

104 The three selected sites for the study are located at Spain (Murcia, Zaragoza and
105 Madrid). They have been chosen because, from the different Wyoming radiosounding data
106 available, they represent different altitudes above sea level.

107 The site situated at lesser height above the sea level is Murcia (62 meters). The
108 radiosoundings are launched at the Territorial Center of AEMET of Murcia, located at 3.5 Km
109 from the city (38° N, 1°9' W).

110 Another region analyzed is focused on the airport of Zaragoza (41°39' N, 1°1' W) at 264
111 meters of altitude above sea level, and 10 km west from the city.

112 Finally, as in the case of Zaragoza, in the Madrid area the radiosoundings are launched
113 at the Airport Adolfo Suarez – Barajas (40°30' N, 3°34' W), located at 611 meters above sea
114 level and 12 Km from the city, in north-east direction.

115 **2.2 Atmospheric Profiles**

116 The spatial and temporal representativeness of radiosounding data is a questionable
117 point in some studies. It is used to consider radiosounding data with a spatial
118 representativeness of 250 km (Huntrieser et al., 1996) and temporal representativeness of 12
119 hours. But these values could be lower in some cases due to the variability of the atmosphere,
120 as happens in mountain areas as Ebro valley (Merino et al., 2013). Different atmospheric
121 profiles are proposed to analyze in these study.

122 *2.2.1 NCEP*

123 The on-line atmospheric correction tool (<http://atmcorr.gsfc.nasa.gov>, last access
124 February 2015) provides the atmospheric parameters needed to make the atmospheric
125 correction in the thermal band of the satellites Landsat-5, Landsat-7 and Landsat-8. The NCEP
126 provides $1^\circ \times 1^\circ$ representative spatial resolution atmospheric profiles every 6 h, as result of
127 four dimensional meteorological data assimilation, including radiosonde (different from WYO
128 soundings), ground and satellite sounder measurements (Barsi et al, 2005). These profiles are
129 processed at 28 different pressure, air temperature and relative humidity, independently of
130 the surface elevation. The corresponding NCEP outputs of temperature ($^\circ\text{C}$) and relative
131 humidity (%), have assigned errors of ± 2 K and $\pm 10\%$, respectively.

132 2.2.2 WYO radiosounding data

133 The Atmospheric Science Department of the University of Wyoming (Laramie, WY,
134 USA) has a database of soundings launched every day at 00:00 and 12:00 UTC at different
135 stations distributed around the world (<http://weather.uwyo.edu/upperair/sounding.html>, last
136 access February 2015).

137 From this database, vertical profiles of height (km), pressure (hPa), temperature ($^\circ\text{C}$),
138 relative humidity (%) and mixing ratio (g/kg), provided at 64 levels, were downloaded for the
139 selected site between 2010 and 2013.

140 These measured variables have uncertainties of ± 1 K for the air temperature and ± 10
141 % for relative humidity.

142 2.2.3 MODIS

143 MODIS (Moderate Resolutions Imaging Spectroradiometer), aboard Terra platform,
144 operates in 36 spectral bands between 0.645 and 14.235 μm , at 705 km of the Earth's surface.
145 Because the characteristics of its spectral bands, MODIS can be used to generate profiles of
146 temperature, relative humidity and dew point at 20 levels of pre-established pressures.
147 Estimations of W are also provided in the atmospheric product named MOD07 (Borbas et al.,

148 2011). Attributed errors to atmospheric parameters provided by MOD07 product are: ± 1.9 °C
149 to the temperature, ± 4 °C to the dew point, ± 10 % to relative Humidity (Borbas et al., 2011).

150 **2.3 Filtering and processing of atmospheric profiles**

151 For a period from 2010 to 2013, a total of 86 scenes for Murcia, 94 for Zaragoza and
152 164 for Madrid have been introduced in the MODTRAN5.2.1 (Berk et al., 2011) RTC to carry
153 out the study. To pick the selected scenes, some restrictions have been taken into account.

154 First we consider as a valid time those days where the satellite passed within 30
155 minutes of the balloon launching. After this first filtering, free of clouds MOD07 scenes over
156 our study region were selected by using the byte codification offered by the MOD07 product,
157 which qualify each pixel as cloudless or not. In a range of 5x5 pixels (Borbas et al., 2011), we
158 require a minimum of 20 cloudless pixels.

159 Finally, only those days for which the three different profiles were available were
160 selected for this study, obtaining the results that are compared in the next section.

161 The atmospheric parameters obtained from running MODTRAN5 are presented in the
162 corresponding spectral range to the MODIS bands 29 (8.55 μm), 31 (11 μm) and 32 (12 μm).

163 **3. RESULTS & DISCUSSIONS**

164 **3.1 Sensitivity analysis**

165 The uncertainty for each of the retrieved parameters is calculated from the
166 uncertainties associated to the different parameters of the profiles commented in section 2.2.
167 The process to calculate the uncertainty associated to the different parameters that
168 characterize the atmosphere is explained. First, MODTRAN is run with the original profiles,
169 then, MODTRAN is run again with profiles that include the original profiles plus the uncertainty
170 associated to their parameters. Finally, the uncertainty associated to each parameter τ_i ,

171 $L_{i,hem}^\downarrow, L_i^\uparrow(\theta)$ and W is obtained calculating the difference of the parameters retrieved from
172 the original profiles and the parameters retrieved from the profiles with the associated error.

173 Table 1 shows some statistics rescued from the uncertainties calculated for the
174 atmospheric parameters $\tau_i, L_{i,hem}^\downarrow, L_i^\uparrow(\theta)$ of the three sources. These values are calculated
175 taking all data available per band and parameter, not dividing the data for each site.

176 ***Insert Table 1 here***

177 The average uncertainty associated to each parameter is lesser for the parameters
178 obtained from MOD07 product, with average values lesser than $\pm 0.5 \text{ Wm}^{-2}\mu\text{m}^{-1}\text{sr}^{-1}$ for $L_i^\uparrow(\theta)$,
179 $\pm 0.6 \text{ Wm}^{-2}\mu\text{m}^{-1}\text{sr}^{-1}$ for $L_{i,hem}^\downarrow$ and ± 0.07 for τ_i . For these cases, the 75% of the uncertainties are
180 under values of $\pm 0.7 \text{ Wm}^{-2}\mu\text{m}^{-1}\text{sr}^{-1}$ for $L_i^\uparrow(\theta)$, $\pm 0.8 \text{ Wm}^{-2}\mu\text{m}^{-1}\text{sr}^{-1}$ for $L_{i,hem}^\downarrow$ and ± 0.09 for τ_i . In
181 the case of NCEP parameters, the uncertainties are higher, with average values lesser than
182 $\pm 0.8 \text{ Wm}^{-2}\mu\text{m}^{-1}\text{sr}^{-1}$, $\pm 1.1 \text{ Wm}^{-2}\mu\text{m}^{-1}\text{sr}^{-1}$ and ± 0.11 for $L_i^\uparrow(\theta)$, $L_{i,hem}^\downarrow$ and τ_i , respectively. For
183 these cases, the 75% of the uncertainties are minor than $\pm 1.13 \text{ Wm}^{-2}\mu\text{m}^{-1}\text{sr}^{-1}$, $\pm 1.5 \text{ Wm}^{-2}\mu\text{m}^{-1}\text{sr}^{-1}$
184 ¹ and ± 0.14 for $L_i^\uparrow(\theta)$, $L_{i,hem}^\downarrow$ and τ_i , respectively. And for WYO parameters, average
185 uncertainties are lesser than $\pm 0.6 \text{ Wm}^{-2}\mu\text{m}^{-1}\text{sr}^{-1}$, $\pm 0.8 \text{ Wm}^{-2}\mu\text{m}^{-1}\text{sr}^{-1}$ and ± 0.08 for $L_i^\uparrow(\theta)$, $L_{i,hem}^\downarrow$
186 and τ_i , respectively. For these cases, the 75% of the uncertainties are down to $\pm 0.8 \text{ Wm}^{-2}\mu\text{m}^{-1}\text{sr}^{-1}$,
187 $\pm 1.1 \text{ Wm}^{-2}\mu\text{m}^{-1}\text{sr}^{-1}$ and ± 0.11 for $L_i^\uparrow(\theta)$, $L_{i,hem}^\downarrow$ and τ_i , respectively.

188 All this uncertainties values correspond to band 32 of each parameter and source,
189 since it is the band with greater associated uncertainties. Moreover, the average values are in
190 most cases equal or very close to the median (Q_2).

191 The results of uncertainties obtained for W are shown in table 2.

192 ***Insert Table 2 here***

193 As happens with the other parameters, W retrieved from NCEP profiles provide the
194 greater uncertainties, with an average of ± 0.7 cm, while for MOD07 and WYO they are of 0.23
195 cm and 0.55 cm. In these cases, the 75% of the parameters have an associated uncertainty
196 lower than ± 0.9 cm, ± 0.3 cm and 0.7 cm for NCEP, MOD07 and WYO, respectively.

197 **3.2 Total column water vapor content**

198 The quantity of water vapor at the atmosphere in a column per unit area (W) is a
199 factor not used directly in the RTE, but with a key importance in the transmissivity, the
200 upwelling and downwelling radiances retrievals. For instance, in García-Santos et al. (2012) the
201 W is obtained through a relationship of downwelling atmospheric radiances, measured at the
202 TIR region in two different angles (0° and 55°).

203 Therefore, when no radiosounding data are available from a launched balloon, it is
204 important to assure that the selected alternative atmospheric profile is able to characterize
205 the atmosphere by means of a realistic W .

206 *Insert Figure 1 here*

207 *Insert Table 3 here*

208 Figure 1 shows the comparison of the W values from NCEP and MOD07 profiles with
209 those from WYO soundings for the full dataset of each site. Comparing results of figure 1 and
210 table 3, it can be observed that W_{NCEP} fit better with W_{WYO} than W_{MOD07} values for the three
211 areas, with an average correlation of 0.93, a BIAS of 0.02 cm and a RMSE of ± 0.2 cm for
212 Zaragoza and a BIAS of 0.2 cm and a RMSE of ± 0.3 cm for Murcia and Madrid, while W_{MOD07}
213 and W_{WYO} present an average correlation of 0.76, a BIAS of -0.05 cm and a RMSE of ± 0.4 cm
214 for Zaragoza, a BIAS of +0.09 and a RMSE of ± 0.3 cm for Murcia and a BIAS of -0.11 and a
215 RMSE of ± 0.3 cm for Madrid. Note that atmospheric MOD07 profiles show a significant
216 underestimation in the W retrievals respect to W_{WYO} values for wet atmospheres ($W_{\text{WYO}} \geq 2.5$

217 cm) on the sites of Zaragoza and Madrid while NCEP profiles retrieve well-correlated W values
218 even for very humid atmospheres ($W_{WYO} \geq 3$ cm), although for Murcia they are overestimated.

219 **3.3 Atmospheric transmissivity**

220 The atmospheric transmissivity (τ) retrieved from the NCEP and MOD07 profiles are
221 also compared with the corresponding values, calculated from radiosounding of the University
222 of Wyoming (τ_{WYO}). Figure 2 shows this comparison between the three selected profiles for the
223 three MODIS thermal bands 29 (8.55 μm), 31 (11 μm) and 32 (12 μm).

224 *Insert Figure 2 here*

225 *Insert Table 4 here*

226 Results of τ comparison depicted in Figure 2 and statistics in table 4, show clearly again
227 a better agreement between NCEP profiles and WYO data than MOD07 profiles. Average
228 regression coefficients for NCEP are 0.93 for band 29, and 0.91 for bands 31 and 32. For
229 Murcia and Madrid, τ is underestimated, with BIAS among -0.013 and -0.019, depending on
230 the MODIS band analyzed, while for Zaragoza it is overestimated, with BIAS among 0.006 and
231 0.007. RMSE values are among ± 0.02 and ± 0.04 . Corresponding statistics for MOD07 are:
232 average regression coefficient of 0.71, 0.69 and 0.70 for bands 29, 31 and 32, respectively. In
233 the case of Murcia, τ retrieved is underestimated, with BIAS among -0.07 and -0.05, while for
234 Zaragoza and Madrid it is overestimated, with BIAS values among 0.001 and 0.06. RMSE values
235 for MOD07 are among 0.03 and 0.08. Note that for MOD07 values of BIAS and RMSE values are
236 close to zero as the altitude of the meteorological station is higher. However, for NCEP they
237 are similar for the three areas.

238 NCEP reproduces transmissivity values closer to those obtained from radiosounding
239 data. This is probably due to the good agreement observed between the water vapor retrieved
240 with NCEP and that calculated from WYO soundings data, since W is indirectly included in the

241 transmissivity exponential expression, through W concentration factor (ρ) and the water
242 spectral absorption coefficient (k, see Eq. (4) in García-Santos et al., 2010).

243 **3.4 Upwelling path radiance**

244 The upwelling radiance ($L_{\lambda}^{\uparrow}(\theta)$ or L_{up}), retrieved from NCEP and MOD07 atmospheric
245 profiles, is compared with that from the radiosoundings of University of Wyoming ($L_{up_{WYO}}$).
246 The results are shown in figure 3 for the three thermal bands of sensor MODIS.

247 *Insert Figure 3 here*

248 *Insert Table 5 here*

249 Graphics of figure 3 show a better agreement for NCEP results than for MOD07, when
250 comparing with L_{up} calculated from WYO profiles. From statistics in table 5, the average
251 regression coefficient for NCEP is 0.95 for band 29 and 0.92 for bands 31 and 32. For Murcia
252 and Madrid, BIAS values are among $0.08 \text{ Wm}^{-2}\mu\text{m}^{-1}\text{sr}^{-1}$ and $0.15 \text{ Wm}^{-2}\mu\text{m}^{-1}\text{sr}^{-1}$, and RMSE
253 among $\pm 0.13 \text{ Wm}^{-2}\mu\text{m}^{-1}\text{sr}^{-1}$ and $\pm 0.3 \text{ Wm}^{-2}\mu\text{m}^{-1}\text{sr}^{-1}$. For Zaragoza, BIAS is of $-0.06 \text{ Wm}^{-2}\mu\text{m}^{-1}\text{sr}^{-1}$
254 and $-0.07 \text{ Wm}^{-2}\mu\text{m}^{-1}\text{sr}^{-1}$, depending of the spectral band, and RMSE among $\pm 0.16 \text{ Wm}^{-2}\mu\text{m}^{-1}\text{sr}^{-1}$
255 and $\pm 0.3 \text{ Wm}^{-2}\mu\text{m}^{-1}\text{sr}^{-1}$. Considering MOD07, the average regression coefficients for each band
256 are: 0.78, 0.73 and 0.73 for bands 29, 31 and 32, respectively. For Murcia, BIAS is among 0.27
257 $\text{Wm}^{-2}\mu\text{m}^{-1}\text{sr}^{-1}$ and $0.35 \text{ Wm}^{-2}\mu\text{m}^{-1}\text{sr}^{-1}$, and RMSE among $\pm 0.4 \text{ Wm}^{-2}\mu\text{m}^{-1}\text{sr}^{-1}$ and $\pm 0.6 \text{ Wm}^{-2}\mu\text{m}^{-1}$
258 sr^{-1} . In the cases of Zaragoza and Madrid, BIAS is among $-0.16 \text{ Wm}^{-2}\mu\text{m}^{-1}\text{sr}^{-1}$ and -0.49 Wm^{-2}
259 $\mu\text{m}^{-1}\text{sr}^{-1}$, and RMSE among $\pm 0.3 \text{ Wm}^{-2}\mu\text{m}^{-1}\text{sr}^{-1}$ and $\pm 0.6 \text{ Wm}^{-2}\mu\text{m}^{-1}\text{sr}^{-1}$.

260 In this case, NCEP data seems to fit better to the curve 1:1. This might be linked to the
261 better agreement in terms of transmissivity shown above. Moreover, it is noticed that the BIAS
262 and RMSE values obtained from both sources are better for high altitudes.

284 **4. Simulation study**

285 In order to analyze the uncertainties in terms of temperature retrieved using the
286 atmospheric parameters calculated in this paper, through different sources, a simulation study
287 was carried out using the SC method (Vlassova et al., 2014). The procedure is as follows:

288 First, for a defined T (in this study three different temperatures were selected: 273 K,
289 293 K and 313 K) a $B_i(T)$ is calculated and a L_i^{TOA} retrieved through Eq. (1) at the three MODIS
290 bands 29, 31 and 32. Atmospheric parameters from the WYO profiles are used to obtain L_i^{TOA}
291 and the surface emissivity (ϵ_i) used in this case, is the corresponding to a rice crop, which has
292 well characterized value of 0.983 ± 0.005 (Coll et al., 2014). Moreover, this value is almost
293 constant in the thermal region.

294 In a second step, a new $B_i(T)$ is calculated through Eq. (1), using the corresponding
295 atmospheric variables obtained from NCEP or MOD07 profiles and the previous calculated
296 L_i^{TOA} . This new $B_i(T)$ is inverted and a new T retrieved. Table 7 shows the average BIAS
297 between original T and those temperatures (for the three selected temperatures) obtained
298 from NCEP and MOD07 profiles and RMSE calculated from all the cases analyzed, at the three
299 MODIS thermal bands.

300 ***Insert Table 7 here***

301 Average results for these three temperatures show that the lowest error are obtained
302 for bands 29 (8.55 μm) and 31 (11 μm). Taking average results of the three sites per MODIS
303 band, band 29 shows a BIAS of +1.3 K and RMSE of ± 0.6 K for NCEP and 1.2 K and ± 1.9 K for
304 MOD07. For band 31, NCEP shows a BIAS of -0.04 K and RMSE of ± 0.6 K and for MOD07 the
305 study shows a BIAS of +0.6 K and RMSE of ± 1.5 K. The greatest error is obtained for band 32
306 (12 μm) where the NCEP profiles show a BIAS of -0.05 K and RMSE of ± 0.8 K, and BIAS of 0.7 K
307 and RMSE of ± 2 K for MOD07, this is reasonable since this band is the most sensitive to water
308 vapor.

309 In summary, the well characterization obtained for the atmospheric parameters τ_i ,
310 $L_{i,hem}^\downarrow$, $L_i^\uparrow(\theta)$, especially for bands 29 and 31, affects directly to the LST retrieval, obtaining
311 lower uncertainties in these cases. Moreover, for MOD07 band 31 obtains better results for
312 LST since it is the less affected band by the water vapor. It is noticed that RMSE obtained are
313 lower in the cases from Madrid, where the altitude is higher and the atmosphere might be
314 better characterized, particularly in profile got from satellite data. These results are in
315 agreement with Coll et al. (2012) for which band 31 has BIAS of -0.3 K and RMSE of ± 0.6 K for
316 NCEP and BIAS of -0.5 K and RMSE of ± 0.8 K for MOD07, and Li et al. (2013) with BIAS of 0.37 K
317 and RMSE of ± 1.16 K for NCEP using the IRS thermal band (10.5 -12.5 μm) and BIAS of 0.20 K
318 and RMSE of ± 1.2 K for MOD07, concluding that NCEP offers less error than MOD07 profiles.

319

320 5. CONCLUSIONS

321 This work has analyzed the differences appeared when characterizing the atmosphere
322 from different atmospheric profiles: NCEP, MOD07_L2 product and radiosoundings from the
323 University of Wyoming (used as reference).

324 Atmospheric parameters in the TIR region 8-14 μm (upwelling radiance, downwelling
325 hemispheric radiance and atmospheric transmissivity) were obtained for three thermal bands
326 (29, 31 and 32) of sensor MODIS.

327 It is concluded that both NCEP and MOD07 profiles show an acceptable agreement
328 with results from WYO profiles. However, despite NCEP profiles show greater uncertainties
329 estimating the atmospheric parameter, because their spatial and temporal interpolation
330 procedure, results show a better fit with the reference WYO values at the three spectral bands.
331 Conversely, the values obtained from MOD07 show lower uncertainties than the NCEP, but
332 their results are more scattered and less correlated with WYO results. So water vapour content

333 is usually underestimated for MOD07, in particular for wet atmospheres ($w > 2.5$ cm) and high
334 altitudes. This underestimation is propagated to the others parameters, underestimating in the
335 same way upwelling and downwelling radiance and overestimating the atmospheric
336 transmissivity.

337 A simulation study was carried out to analyze the effect of each atmospheric profile
338 retrieving the surface temperature when applying the SC method. Results showed RMSE
339 between ± 0.6 - ± 0.9 K for NCEP and between ± 1.3 - ± 3 K for MOD07, depending on the spectral
340 band and the altitude of the study site. So for bands 29 and 31, temperature uncertainties
341 show the best results, probably because band 32 is more affected by the water vapor.
342 Moreover, it is observed better results on the LST retrievals for high altitudes, leading to lower
343 errors in the way that the altitude is greater. This is probably due to MOD07 product get less
344 data for lower altitudes. For example, in the case of Murcia, the first data level is around 120
345 meters, 60 meters over the ground altitude. While for higher altitudes, that measure is more
346 accurate.

347 Radiosounding data is probably the best way to characterize the atmosphere, but
348 rarely there are radiosoundings available at the time and location of interest. This study has
349 shown that using as alternative NCEP profiles leads to optimal results when characterizing the
350 atmosphere. In addition, good results are obtained for MODIS band 29 ($8.55 \mu\text{m}$), which has
351 not been studied previously in other works, with uncertainties similar to those obtained for
352 MODIS band 31 ($11 \mu\text{m}$).

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407 Tables

408 Table 1. Statistics of the uncertainties obtained for the different atmospheric parameters (L_{up} , L_{down} , τ) in each MODIS spectral band and for the three
 409 sources. Max and min represents the maximum and minor value obtained, av is the average uncertainty and Q1, Q2, Q3 are the first, second and third
 410 quartile, respectively.

	NCEP						MOD07						WYO					
	BAND 29						BAND 29						BAND 29					
	max	min	av.	Q1	Q2	Q3	max	min	av.	Q1	Q2	Q3	max	min	av.	Q1	Q2	Q3
L_{up} ($Wm^{-2}\mu m^{-1}sr^{-1}$)	1.1	0.13	0.5	0.3	0.5	0.7	0.8	0.08	0.3	0.2	0.5	0.4	0.7	0.07	0.4	0.2	0.3	0.5
L_{down} ($Wm^{-2}\mu m^{-1}sr^{-1}$)	1.5	0.2	0.8	0.5	0.8	1.0	1.0	0.11	0.5	0.3	0.8	0.6	1.1	0.12	0.5	0.3	0.5	0.7
τ	0.11	0.02	0.06	0.05	0.06	0.08	0.09	0.01	0.04	0.03	0.06	0.05	0.09	0.01	0.05	0.04	0.05	0.06
	BAND 31						BAND 31						BAND 31					
L_{up} ($Wm^{-2}\mu m^{-1}sr^{-1}$)	1.7	0.09	0.7	0.4	0.7	1.0	1.3	0.05	0.5	0.2	0.4	0.6	1.2	0.06	0.5	0.3	0.5	0.7
L_{down} ($Wm^{-2}\mu m^{-1}sr^{-1}$)	2.0	0.15	1.0	0.6	1.1	1.4	1.4	0.08	0.6	0.3	0.5	0.7	1.6	0.10	0.7	0.4	0.7	1.0
τ	0.17	0.01	0.09	0.05	0.09	0.12	0.14	0.01	0.06	0.03	0.05	0.08	0.13	0.01	0.06	0.04	0.06	0.09
	BAND 32						BAND 32						BAND 32					
L_{up} ($Wm^{-2}\mu m^{-1}sr^{-1}$)	1.7	0.12	0.8	0.5	0.9	1.1	1.3	0.07	0.5	0.3	0.5	0.7	1.3	0.08	0.6	0.3	0.6	0.8
L_{down} ($Wm^{-2}\mu m^{-1}sr^{-1}$)	2.1	0.2	1.13	0.8	1.2	1.5	1.4	0.10	0.6	0.4	0.6	0.8	1.6	0.13	0.8	0.5	0.8	1.1
τ	0.33	0.02	0.11	0.07	0.12	0.14	0.15	0.01	0.07	0.04	0.07	0.09	0.16	0.01	0.08	0.05	0.08	0.11

411

412 **Table 2. Statistics of the uncertainties obtained for the water vapour content each source.**

413 **Max and min represents the maximum and minor value obtained, av is the average**

414 **uncertainty and Q1, Q2, Q3 are the first, second and third quartile, respectively.**

	W (cm)					
	max	min	av	Q1	Q2	Q3
NCEP	1.4	0.19	0.7	0.5	0.7	0.9
MOD07	0.7	0.004	0.2	0.10	0.19	0.3
WYO	1.1	0.13	0.6	0.4	0.5	0.7

415

416 **Table 3. Linear regression parameters obtained from the comparison of W retrieved**

417 **between the NCEP-WYO and MOD07-WYO (figure 1). From the fitting, m represents the**

418 **slope, n is the offset, r^2 is the regression coefficient and σ the standard deviation. In**

419 **addition, it is shown the average BIAS and RMSE of the depicted points.**

NCEP - WYO	m	n	r^2	σ	BIAS	RMSE
Murcia	1.07	0.08	0.93	0.2	0.2	0.3
Zaragoza (cm)	0.93	0.13	0.92	0.2	0.018	0.2
Madrid	0.99	0.22	0.94	0.15	0.2	0.3
MOD07 - WYO	m	n	r^2	σ	BIAS	RMSE
Murcia	0.99	0.11	0.80	0.3	0.09	0.3
Zaragoza (cm)	0.80	0.26	0.74	0.4	-0.05	0.4
Madrid	0.76	0.20	0.74	0.3	-0.11	0.3

420

421

422 **Table 4. Linear regression parameters obtained from the comparison of transmissivity retrieved for three MODIS TIR bands between the NCEP-WYO and**
 423 **MOD07-WYO (figure 2), for three different sites. From the fitting, m represents the slope, n is the offset, r^2 is the regression coefficient and σ the**
 424 **standard deviation. In addition, it is show the average BIAS and RMSE of the depicted points.**

	NCEP						MOD07					
	MURCIA						MURCIA					
	m	n	r^2	σ	BIAS	RMSE	m	n	r^2	σ	BIAS	RMSE
τ_{29}	1.01	-0.02	0.93	0.02	-0.013	0.02	0.98	-0.05	0.74	0.04	-0.07	0.08
τ_{31}	1.04	-0.05	0.91	0.03	-0.016	0.03	0.99	-0.04	0.74	0.05	-0.05	0.07
τ_{32}	1.04	-0.05	0.91	0.03	-0.019	0.04	0.96	-0.03	0.74	0.06	-0.06	0.08
	ZARAGOZA						ZARAGOZA					
τ_{29}	0.92	0.06	0.92	0.02	0.006	0.02	0.71	0.25	0.76	0.03	0.04	0.05
τ_{31}	0.90	0.09	0.89	0.03	0.007	0.03	0.61	0.37	0.72	0.03	0.05	0.06
τ_{32}	0.91	0.08	0.89	0.04	0.009	0.04	0.63	0.35	0.73	0.04	0.06	0.08
	MADRID						MADRID					
τ_{29}	0.94	0.03	0.95	0.013	-0.015	0.02	0.71	0.22	0.63	0.03	0.0014	0.03
τ_{31}	0.95	0.03	0.93	0.017	-0.013	0.02	0.61	0.35	0.62	0.03	0.02	0.04
τ_{32}	0.94	0.03	0.93	0.02	-0.016	0.03	0.62	0.34	0.63	0.04	0.03	0.05

425 Table 5. Linear regression parameters obtained from the comparison of upwelling radiance retrieved for three MODIS TIR bands between the NCEP-WYO
 426 and MOD07-WYO (figure 2), for three different sites. From the fitting, m represents the slope, n is the offset, r^2 is the regression coefficient and σ the
 427 standard deviation. In addition, it is show the average BIAS and RMSE of the depicted points.

		NCEP						MOD07						
		MURCIA						MURCIA						
		m	n	r^2	σ	BIAS	RMSE	m	n	r^2	σ	BIAS	RMSE	
Lup₂₉	(Wm⁻²μm⁻¹sr⁻¹)	1.03	0.03	0.95	0.15	0.08	0.17	1.02	0.24	0.85	0.3	0.3	0.4	
		Lup₃₁	1.05	0.05	0.92	0.2	0.13	0.3	1.01	0.31	0.80	0.4	0.3	0.5
		Lup₃₂	1.04	0.07	0.92	0.3	0.15	0.3	0.98	0.39	0.79	0.4	0.4	0.6
		ZARAGOZA						ZARAGOZA						
Lup₂₉	(Wm⁻²μm⁻¹sr⁻¹)	0.93	0.05	0.94	0.15	-0.06	0.16	0.72	0.13	0.80	0.2	-0.3	0.4	
		Lup₃₁	0.91	0.06	0.90	0.2	-0.06	0.2	0.61	0.12	0.74	0.3	-0.4	0.5
		Lup₃₂	0.91	0.08	0.90	0.3	-0.07	0.3	0.62	0.13	0.75	0.3	-0.5	0.6
		MADRID						MADRID						
Lup₂₉	(Wm⁻²μm⁻¹sr⁻¹)	1.00	0.11	0.97	0.10	0.11	0.15	0.70	0.26	0.70	0.2	-0.16	0.3	
		Lup₃₁	0.98	0.13	0.94	0.14	0.11	0.18	0.59	0.22	0.65	0.2	-0.2	0.3
		Lup₃₂	0.97	0.17	0.94	0.16	0.13	0.13	0.59	0.24	0.65	0.3	-0.3	0.4

428

429

430

431 Table 6. Linear regression parameters obtained from the comparison of downwelling hemispheric radiance retrieved for three MODIS TIR bands between
 432 the NCEP-WYO and MOD07-WYO (figure 2), for three different sites. From the fitting, m represents the slope, n is the offset, r^2 is the regression
 433 coefficient and σ the standard deviation. In addition, it is show the average BIAS and RMSE of the depicted points.

	NCEP						MOD07					
	MURCIA						MURCIA					
	m	n	r^2	σ	BIAS	RMSE	m	n	r^2	σ	BIAS	RMSE
Ldown₂₉	1.08	0.00	0.94	0.26	0.2	0.3	0.90	0.02	0.89	0.30	-0.3	0.4
Ldown₃₁ ($Wm^{-2}\mu m^{-1}sr^{-1}$)	1.06	0.07	0.92	0.34	0.2	0.4	0.78	0.19	0.83	0.37	-0.3	0.5
Ldown₃₂	1.05	0.09	0.92	0.37	0.2	0.4	0.79	0.22	0.83	0.43	-0.3	0.5
	ZARAGOZA						ZARAGOZA					
Ldown₂₉	0.98	0.05	0.94	0.23	-0.014	0.2	0.78	0.15	0.81	0.36	-0.4	0.5
Ldown₃₁ ($Wm^{-2}\mu m^{-1}sr^{-1}$)	0.93	0.08	0.91	0.32	-0.06	0.3	0.66	0.15	0.75	0.40	-0.6	0.7
Ldown₃₂	0.94	0.08	0.91	0.35	-0.07	0.4	0.68	0.16	0.76	0.46	-0.7	0.8
	MADRID						MADRID					
Ldown₂₉	1.04	0.14	0.97	0.15	0.2	0.3	0.65	0.26	0.78	0.28	-0.5	0.6
Ldown₃₁ ($Wm^{-2}\mu m^{-1}sr^{-1}$)	0.99	0.20	0.95	0.20	0.19	0.3	0.50	0.24	0.70	0.27	-0.6	0.7
Ldown₃₂	0.98	0.26	0.95	0.23	0.2	0.3	0.51	0.26	0.70	0.33	-0.7	0.8

434

435

436 **Table 7. BIAS and RMSE calculated from the comparison between the original LST and those**
 437 **retrieved using NCEP or MOD07 atmospheric parameters for three MODIS bans (29, 31, 32)**
 438 **in each one of the selected sites.**

	MURCIA			
	NCEP		MOD07	
Band	BIAS (K)	RMSE (K)	BIAS (K)	RMSE (K)
29	0.08	0.6	3	3
31	-0.05	0.6	1.2	1.9
32	-0.05	0.8	1.6	3
	ZARAGOZA			
29	0.15	0.6	0.15	1.3
31	0.011	0.7	-0.03	1.4
32	-0.004	0.9	-0.07	2
	MADRID			
29	0.03	0.6	1.3	1.7
31	-0.09	0.6	0.6	1.2
32	-0.09	0.7	0.8	2

439

440 **FIGURE CAPTIONS**

441 **Figure 1. Differences NCEP-WYO, and MOD07-WYO, versus W.**

442 **Figure 2. Linear regression of the comparison between NCEP and WYO (three top graphs)**
443 **and MODIS-WYO (three bottom graphs), in terms of estimated τ for the three MODIS**
444 **thermal bands 29 (8.55 μm), 31 (11 μm) and 32 (12 μm).**

445 **Figure 3. Linear regression of the comparison between NCEP and WYO (three top graphs)**
446 **and MODIS-WYO (three bottom graphs), in terms of estimated L_{up} for the three MODIS**
447 **thermal bands 29 (8.55 μm), 31 (11 μm) and 32 (12 μm).**

448 **Figure 4. Linear regression of the comparison between NCEP and WYO (three top graphs)**
449 **and MODIS-WYO (three bottom graphs), in terms of estimated L_{down} for the three MODIS**
450 **thermal bands 29 (8.55 μm), 31 (11 μm) and 32 (12 μm).**

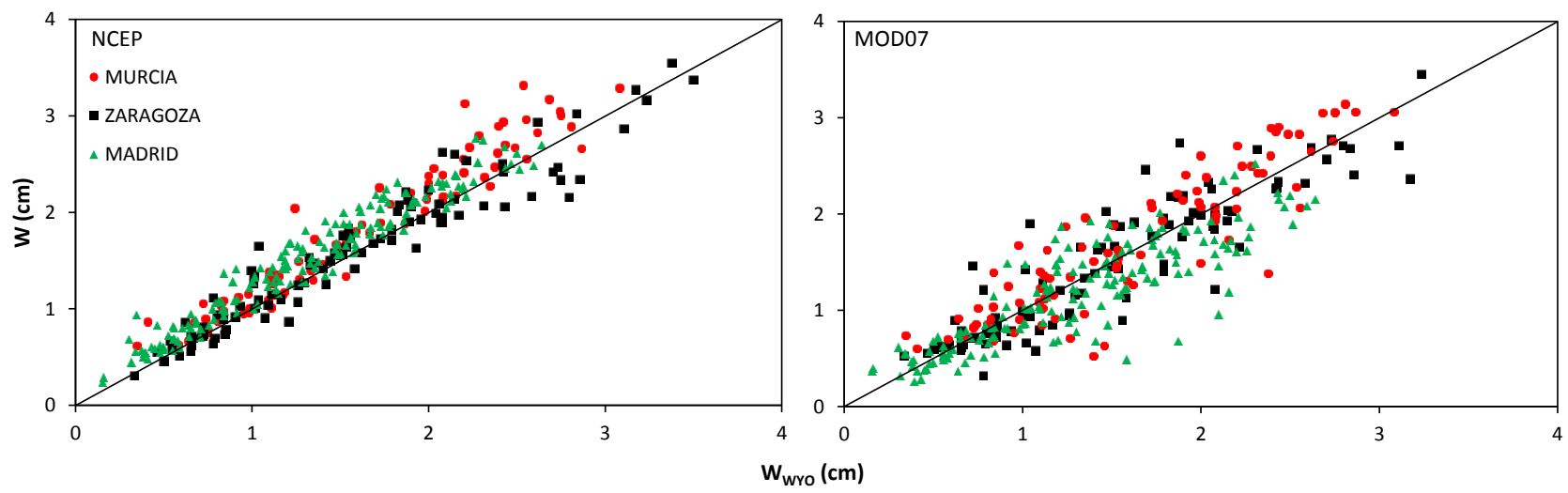


Figure 1

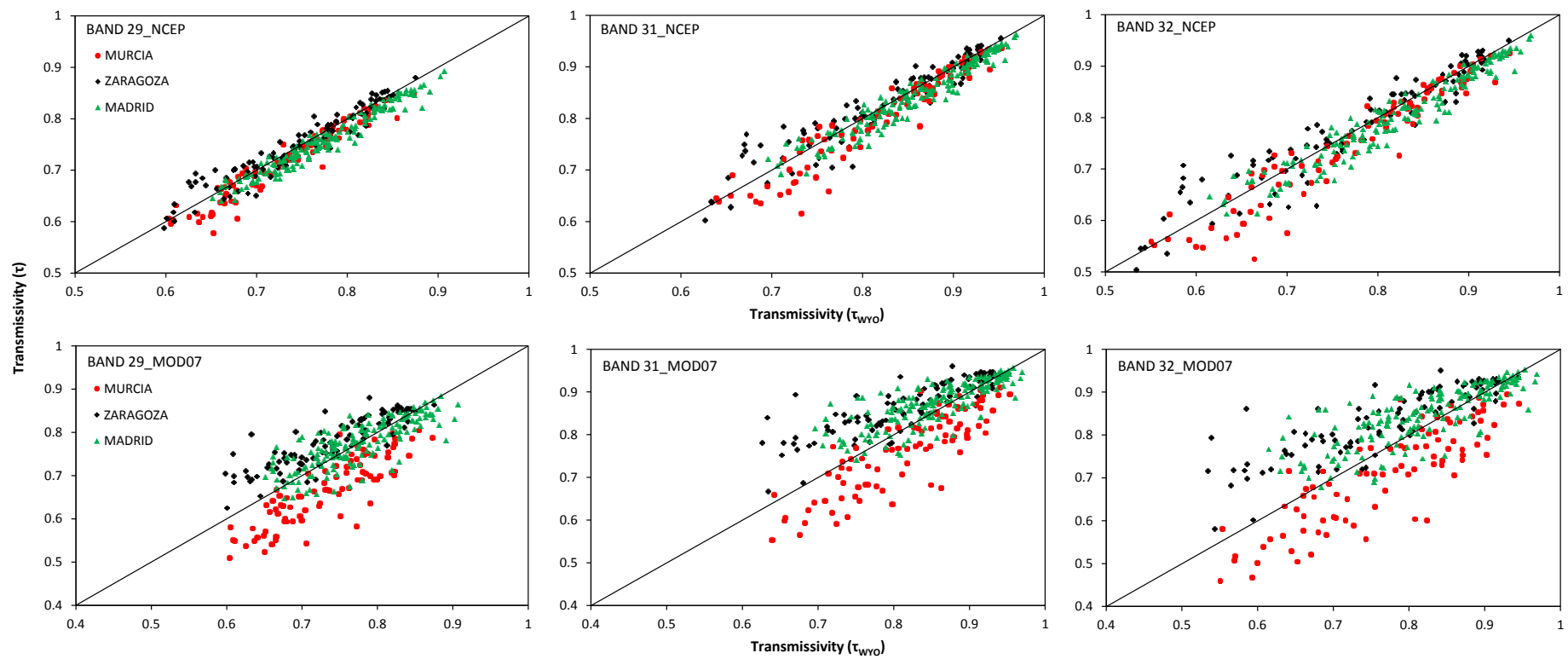


Figure 2

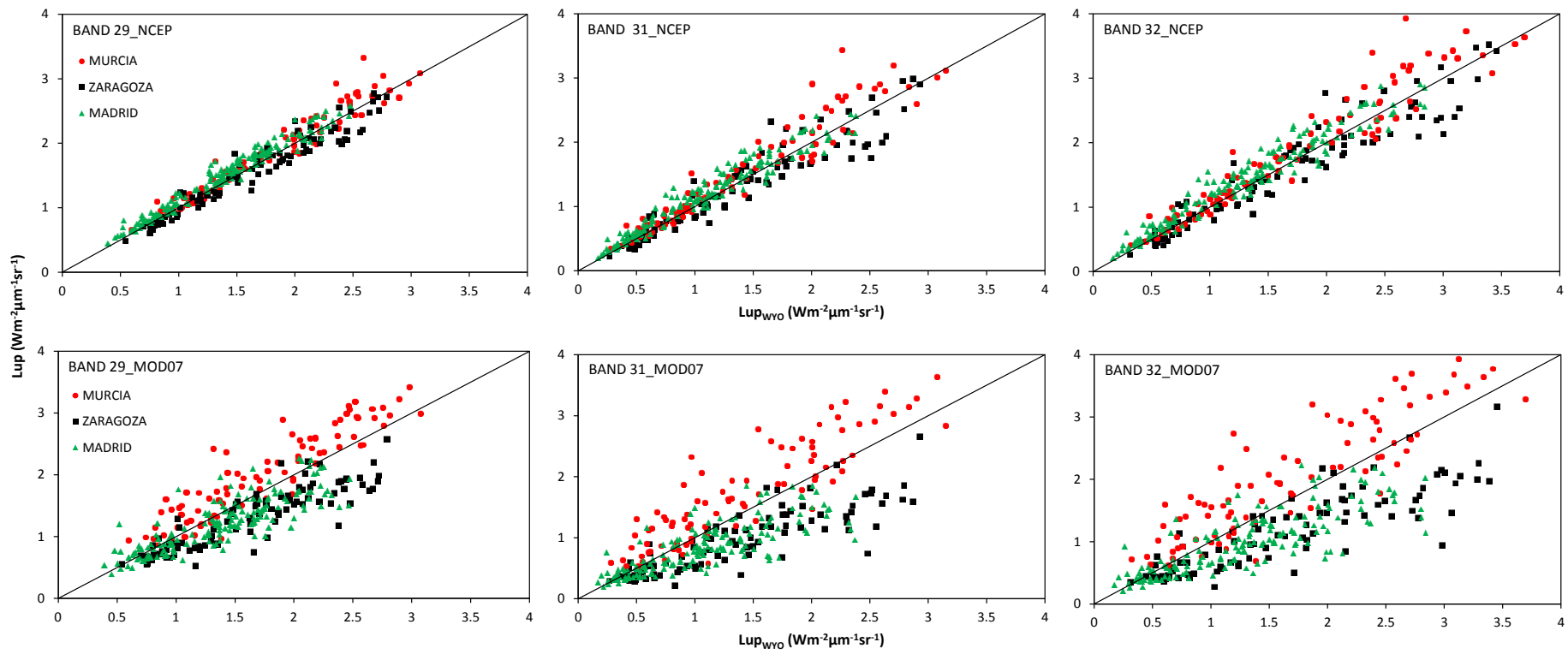


Figure 3

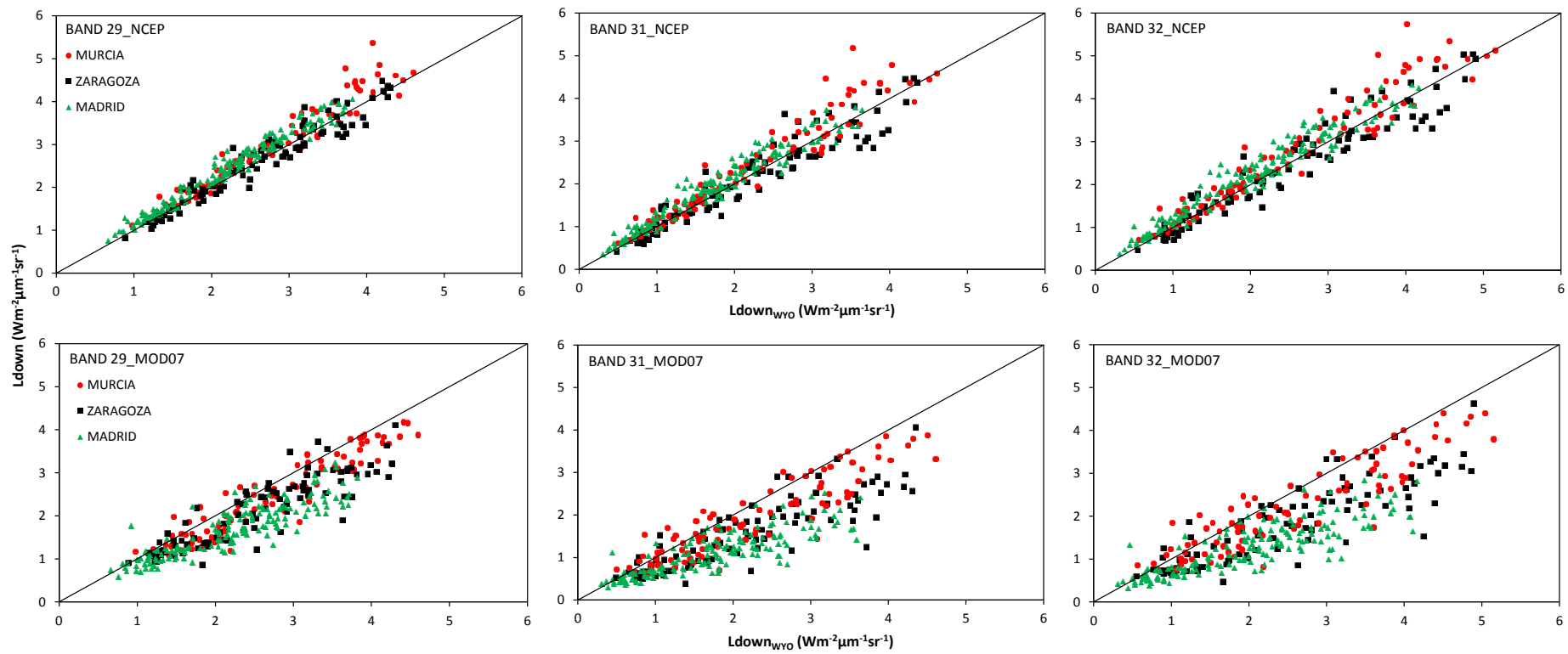


Figure 4